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Low-level jet characteristics over the Arctic Ocean in spring and summer

L. Jakobson[1], T. Vihma[2], E. Jakobson[3,4], T. Palo[1], A. Männik[5], and J. Jaagus[1]

[1]Department of Geography, University of Tartu, Vanemuise 46, 51014, Tartu, Estonia
[2]Finnish Meteorological Institute, P.O. Box 503, 00101, Helsinki, Finland
[3]Tartu Observatory, 61602, Tõravere, Tartumaa, Estonia
[4]Department of Physics, University of Tartu, Tähe 4, 51010, Tartu, Estonia
[5]Estonian Meteorological and Hydrological Institute, Mustamäe tee 33, 10616, Tallinn, Estonia

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Correspondence to: L. Jakobson (liisi.jakobson@ut.ee) and T. Vihma (timo.vihma@fmi.fi)

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Abstract

Low-level jets (LLJ) are important for turbulence in the stably stratified atmospheric boundary layer, but their occurrence, properties, and generation mechanisms in the Arctic are not well known. We analysed LLJs over the central Arctic Ocean in spring and summer 2007 on the bases of data collected in the drifting ice station Tara. Instead of traditional radiosonde soundings, data from tethersonde soundings with a high vertical resolution were used. The Tara results showed a lower occurrence of LLJs (46 %) than many previous studies over polar sea ice. Strong jet core winds contributed to growth of the turbulent layer. Complex relationship between the jet core height and the temperature inversion top height were detected: substantial correlation ($r = 0.72$; $p < 0.01$) occurred when the jet core was above the turbulent layer, but inside the turbulent layer there was no correlation. The most important forcing mechanism for LLJs was baroclinicity, which was responsible for generation of strong and warm LLJs, which on average occurred at lower altitudes than other jets. Baroclinic jets were mostly associated to transient cyclones instead of the climatological air temperature gradients. Besides baroclinicity, cases related to inertial oscillations, gusts, and fronts were detected. In approximately 50 % of the observed LLJs the generation mechanism remained unclear, but in most of these cases the wind speed was strong in the whole vertical profile, the jet core representing only a weak maximum. Further research needs on LLJs in the Arctic include investigation of low-level jet streams and their effects on the sea ice drift and atmospheric moisture transport.

1 Introduction

Numerous recent studies have demonstrated major changes in the climate system of the central Arctic. Air temperatures have increased (e.g. Walsh et al., 2011) and the sea ice melt season has become longer (Maksimovich and Vihma, 2012). Sea ice has become thinner, its drift velocities have increased, and its extent has strongly decreased in summer and autumn (Stroeve et al., 2012). Arctic warming during the 21st century is very likely to exceed the global mean warming but, simultaneously, the scatter between various climate model projections for the 21st century is particularly large in the Arctic (Christensen et al., 2007). Further, climate models have large problems in simulating the recent changes in the Arctic sea ice cover (Stroeve et al., 2007), and even atmospheric reanalyses include major errors over the Arctic sea ice (Jakobson et al., 2012).

Errors in both climate models (Tjernström et al., 2005) and numerical weather prediction models (Atlaskin and Vihma, 2012) tend to be largest in conditions of a stable boundary layer (SBL). There are several reasons that make SBL a challenge for models (Steeneveld et al., 2006; Atlaskin and Vihma, 2012). One of them is related to the low-level jet (LLJ, a low-altitude maximum in the vertical profile of the wind speed), which commonly occurs in conditions of a SBL. In a SBL, turbulence near the Earth surface is weak. Hence, the wind shear below the core of a LLJ may be the main source of turbulence (Mahrt, 2002; Mäkiranta et al., 2011). This results in a top-down structure of the SBL, but the model parameterizations are not designed for such conditions. Further, a LLJ often occurs intermittently, so that the shear-driven turbulence is also intermittent, which is another major challenge for models (Mahrt, 2002; Costa et al., 2011). A LLJ is often detected only as a maximum in the vertical wind profile, without any particular three-dimensional structure. Some LLJs are, however, associated with a narrow horizontal zone of a high-speed flow, called as a low-level jet stream (Stensrud, 1996).

In the Arctic Ocean, LLJs may also affect the motion of the sea ice margin (Langland et al., 1989), which further affects the sea ice mass balance. There are, however, not many detailed studies on the occurrence and generation mechanisms of LLJs over the Arctic sea ice. Nearby the coasts of Greenland and Svalbard, LLJs are often related to katabatic winds (Heinemann, 2004; Vihma et al., 2011) or more complex orographic effects (Samelson and Barbour, 2008; Esau and Repina, 2012). LLJs are also common over sea ice far from orographic influence: Langland et al. (1989) observed LLJs related to an ice breeze – a sea-breeze type mesoscale circulation. Vihma et al. (1998)

observed LLJs over the ice edge zone in the Denmark Strait; the strongest LLJs were generated by baroclinicity. Andreas et al. (2000) observed a high frequency of occurrence (80 %) of LLJs over the Antarctic sea ice zone, and suggested that they were primarily due to inertial oscillations: a mechanism analogous to the classical nocturnal jet (Blackadar, 1957; Thorpe and Guymer, 1977) but in the Antarctic related to synoptic-scale changes in the atmospheric boundary layer (ABL) stratification. Inertial oscillations generated by spatial changes in surface roughness and stratification may also generate LLJs, as observed by Smedman et al. (1993) and Vihma and Brümmer (2002) in the Baltic Sea. Also ReVelle and Nilsson (2008) associated LLJs to inertial oscillations. They observed a LLJ in some 2/3 of all rawinsonde soundings during a three-month-long Arctic Ocean expedition in summer 1996.

Insufficiency of high-resolution data on the vertical profiles of wind speed is the largest impediment for exploring LLJs over the Arctic Ocean. Rawinsonde soundings are only taken during cruises of a few research vessels (Lüpkes et al., 2010; Tjernström et al., 2012), most of the cruises lasting no more than approximately a month in the sea ice zone. Radiosonde soundings have been made during the Russian drifting ice stations since 1950s, but decades ago the data quality and vertical resolution have not been sufficient to yield good statistics of LLJs. This was demonstrated by Andreas et al. (2000), who carried out tethersonde soundings over the Antarctic sea ice zone in 1992 and showed their superior applicability in LLJ studies compared to traditional radiosonde soundings. Tethersonde soundings, were also carried out during the drifting station Surface Heat Budget of the Arctic Ocean (SHEBA, Uttal et al., 2002), but we are not aware of studies on LLJs based on these data.

The next major tethersonde sounding campaign over the Arctic sea ice took place during the drifting ice station Tara in spring and summer 2007 (Gascard et al., 2008; Vihma et al., 2008); the data collected forms the basis of our study. The objective of this paper is to quantify characteristics of LLJs over the Arctic Ocean in spring and summer and to find out their most important formation mechanisms. Some mechanisms that elsewhere generate jets (e.g. terrain effects and the diurnal cycle) are not active over a flat sea ice surface very close to the North Pole. The potential generation mechanisms include baroclinicity, inertial oscillations, fronts, and gusts; these will be studied in detail.

2 Observations and methods

2.1 Field observations

Meteorological observations from the drifting ice station Tara were carried out in the central Arctic Ocean from March to September 2007 (Fig. 1). The tethersonde sounding period lasted from 25 April to 31 August.

A Vaisala DigiCORA Tethersonde System was used to measure the vertical profiles of the wind speed, air temperature, relative humidity, and wind direction (Vihma et al., 2008). In this study we focus on the wind profile, but information on the temperature profile is also used to interpret the wind conditions. Our tethersonde system consisted of a 7 m^3 balloon filled by helium, tetherline, winch, and three sondes with 20 m vertical intervals. Due to the risk of breaking the balloon or tetherline, the measurements were only carried out under wind speeds lower than 15 $m\,s^{-1}$ in the whole profile. The balloon was ascended as high as possible (the average top height of the soundings was 1240 m), and the data were recorded with about 5 m intervals. Though the winch was spooling with constant speed of 1 to 1.5 $m\,s^{-1}$, the balloon did not gain height with a constant speed. The balloon did not rise up straight but drifted along the wind. Hence, the recorded wind speed values were systematically higher during descent than ascent (usually from 0.5 to 2 $m\,s^{-1}$). Hence, an average profile was calculated (for each sensor separately) on the basis of the ascending and descending profiles. This averaging (over every 20 m) yields more reliable results, although some information on temporal variations is lost.

Due to the vicinity of the geomagnetic pole, the observed wind direction was very sensitive to even a small tilt of the compass, which was inevitable in the tethersonde

system (the three vanes often showed different wind directions although the digital compasses were calibrated before every sounding). Hence, the wind direction results from tethersonde system were not used for detailed analysis.

The temporal distribution of soundings was made as equal as possible (depending on the weather conditions and technical possibilities). In consequence of two 24-h-long intensive measurement periods, there were more soundings in July. Also, there was one 24 h measurement period in May and another one in June. In April there were few soundings because the first sounding day was only on 25 April. 57 % of soundings were made between 07:00 and 12:00 LST.

In addition to tethersonde soundings, the air temperature and wind speed were measured at a 10-m-high weather mast (Aanderaa AWS 2700) at the heights of 1, 2, 5 and 10 m, the air relative humidity at 2 m and wind direction at 10 m.

2.2 Jet definitions and analyses

A LLJ was defined following Stull (1988) as the level with a local wind speed maximum of more than 2 $m\,s^{-1}$ greater than wind speeds above it. The level of maximum wind was defined as the jet core (z_j). The difference between z_j and the subsequent wind speed minimum above (z_a) was defined as the jet depth ($z_a - z_j$). The wind speed difference between the core speed (U_j) and the minimum speed above (U_a) was defined as the jet strength ($U_j - U_a$). The level of maximum air temperature was defined as the temperature inversion top (z_t). The difference between the z_t and the previous temperature minimum below (z_b) was defined as the temperature inversion depth ($z_t - z_b$). The air temperature difference between the inversion top temperature (T_t) and the minimum temperature below (T_b) was defined as the temperature inversion strength ($T_t - T_b$). Figure 2 illustrates these variables. In the illustrated example sounding from 10 August 2007, the data allow identifying a LLJ in the wind speed profile with a core speed of 8.4 $m\,s^{-1}$ at the height of 180 m. The wind is remarkably weak near the surface and around 800 m. The top of inversion (230 m) is slightly above the jet core (180 m).

The inversion strength is only 1.6 °C, but the jet strength (5.7 m s^{-1}) is larger than the average observed at Tara.

In the sounding period from 25 April to 31 August, there were a total of 95 soundings in 39 sounding days. For LLJ statistics, one sounding per each day was selected. Such a selection was needed because the LLJs observed were not necessarily independent of each other (up to eight soundings per day were made). To count the occurrence of LLJs, the highest sounding per each day was chosen; 18 of these 39 cases, i.e. 46 %, included a LLJ. To count the other properties of LLJ, the existence of LLJ was the criteria for choosing one sounding per day (one or more LLJs were observed in 25 of the 39 days). From sounding days with more than one LLJ observed, the highest sounding was chosen. All the observed LLJ profiles were applied in analyses of the generation mechanisms of LLJs (43 profiles among 95 soundings). To summarize, we had a total of 95 soundings, 43 LLJs observed, and 25 soundings were included in analyses of LLJ properties.

The bulk-Richardson number (Ri), which is a non-dimensional parameter describing the ratio of buoyancy and wind shear in turbulence production (e.g. Kaimal and Finnigan, 1994), was used to represent static stability. As Andreas et al. (2000), we calculated Ri from the surface to each observation height:

$$Ri(z) = \frac{gz}{\Theta(z)} \frac{\Theta(z) - \Theta_s}{v^2(z)}. \quad (1)$$

Here, g is the acceleration of gravity; z is the observation height; $\Theta(z)$ and $v(z)$ are the potential temperature and wind speed at z; and Θ_s is the potential temperature at the height of 10 m, which was the first averaging height of the tethersonde data (in cases of LLJs, the 10 m temperature was within ±0.3 °C of the 1 m temperature recorded in the weather mast). If $Ri(z)$ was smaller than the critical Richardson number (Ri_{cr}), the layer up to the height $z$was considered to be turbulent. The Ri_{cr} has no unambiguous value; empirically based suggestions in the literature range from 0.2 to 1.0 (Galperin et al.,

2007). We took $Ri_{cr} = 0.4$, similarly to Andreas et al. (2000). The lowest level for which $Ri(z) \geq Ri_{cr}$ is indicated as z_{Ri} and is assumed to be the top of the turbulent layer.

The average wind profile plus/minus standard deviation, based on 43 soundings (all LLJs), is shown in Fig. 3. The mean profile includes a wind maximum at the height of 200 m, but it does not meet the criterion for a LLJ. There is no significant difference in stability between the cases with and without a LLJ (not shown).

2.3 Supporting material and analyses

We applied the operational analyses of the European Centre for Medium-Range Weather Forecasts (ECMWF) to calculate air temperature gradients and the thermal wind, and to detect fronts. We also applied the METEX backward trajectory calculator (http://db.cger.nies.go.jp/metex/trajectory.html), which utilizes the NCEP/NCAR reanalayses, to calculate 72 h backward trajectories for the LLJ cases observed.

3 Generation mechanisms of low-level jets

As mentioned, a low-level jet in the wind profile was detected in 43 of the 95 soundings. All the observed LLJ profiles were applied in analyses of the generation mechanisms. LLJs can be generated by a variety of mechanisms, including (a) baroclinicity, (b) inertial oscillations due to temporal and spatial variations in the surface friction, (c) directional shear of other origin, (d) mesoscale circulations such as an ice breeze, (e) fronts, and (f) gusts.

3.1 Baroclinicity

The baroclinicity related to a horizontal temperature gradient may generate a LLJ at the level above which the decreasing geostrophic wind dominates and below which the effect of surface friction dominates. Air temperature fields based on the ECMWF operational analyses were used to identify the cases with geostrophic wind speed decreasing

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with height. The equations for thermal wind are as follows (e.g. Stull, 2009):

$$\frac{\partial U_g}{\partial z} = -\frac{g}{f_c T}\frac{\partial T}{\partial y}, \quad (2)$$

$$\frac{\partial V_g}{\partial z} = +\frac{g}{f_c T}\frac{\partial T}{\partial x}. \quad (3)$$

Here, U_g is the eastward and V_g the northward component of geostrophic wind; f_c is the Coriolis parameter; T is the temperature; x and y are coordinates towards east and north, respectively. The geostrophic wind speed at the surface and z_a were calculated. If the geostrophic wind speed was at least $2\,\mathrm{m\,s^{-1}}$ smaller atz_a than at surface, the baroclinicity criteria was fulfilled. 13 cases (of 43) fulfilled the criteria; three of these cases were also detected as potentially generated by inertial oscillations (see below).

For baroclinic jets, the mean jet strength ($U_j - U_a$) was $0.9\,\mathrm{m\,s^{-1}}$ larger than for jets which had no baroclinicity forcing mechanism (the difference is significant at the confidence level $p < 0.05$). The mean inversion base temperature (T_b) and inversion top temperature (T_t) were, respectively 6.3 °C and 6.2 °C higher ($p < 0.01$) than for jets which had no baroclinicity forcing mechanism. The mean z_j of baroclinic jets (265 m) occurred 172 m lower than in the case of other jets ($p < 0.01$). The baroclinicity forcing mechanism was more important in July and August (11 cases) than in April–June (two cases).

Comparisons of instantaneous and seasonal mean pressure and temperature fields, based on the ECMWF operational analyses, showed that the baroclinic LLJs were related to transient cyclones. Also the seasonal mean temperature fields at 925 hPa level (altitude where baroclinity typically reduced the geostrophic wind speed) included horizontal gradients in the study region, but these were not large enough or did not have the correct orientation to generate LLJs at the observation site (Fig. 4). If the baroclinic LLJs were due to the seasonal mean temperature field, the trajectories (at least their later parts before reaching Tara) should be aligned parallel to the isotherms

so that cold air lies on the right of the wind vector, but this was the case only once in July–August and never in April–June (Fig. 4).

3.2 Inertial oscillations

Inertial oscillations related to the Coriolis force and ceasing of frictional drag may induce a LLJ later at night (Blackadar, 1957) or after storms, when the stable stratification is re-established (Andreas et al., 2000). To distinguish LLJs generated by inertial oscillations, one could study the history of stratification, as the occurrence of an inertial LLJ should be preceded by a period of neutral or unstable stratification. It is, however, not well known how long inertial oscillations may persist before being damped (i.e. how far in the history we should look at). Further, the neutral or unstable stratification may have occurred far from the sounding site. Hence, we only paid attention to the jet core height. LLJs generated by inertial oscillations typically have their core close to the top of the stable boundary layer (Thorpe and Guymer, 1977; Andreas et al., 2000). As it is not possible to give an exact criteria for the threshold stratification for occurrence of turbulence, we classify as potentially inertial those jets that have their core above the lowest level where $Ri \geq 0.2$ but below the lowest level where $Ri \geq 0.7$. Among our 43 detected LLJs, seven cases fulfilled this criterion. Note that LLJs generated by other forcing mechanisms may also have their core heights in the above-mentioned layer. In fact, three of the seven cases were also detected as baroclinical ones. Jets that were potentially generated by inertial oscillation had $1.5\,\mathrm{m\,s^{-1}}$ higher ($p < 0.05$) wind speed at jet core (U_j) than jets which had no inertial oscillation forcing mechanism.

We also studied the possibility of LLJs generated by inertial oscillations due to a spatial change in surface friction, as observed by Vihma and Brümmer (2002). The travel time of air mass between the sea ice margin and Tara was calculated. Theoretically, the maximum LLJ occurs after slightly less than half of the inertial period (Blackadar, 1957), and the next maxima occur during the second, third, and fourth oscillations, i.e. after approximately 1.4, 2.4, and 3.4 times the inertial period. There were four cases of LLJs when the air mass flew from the open Fram Strait, thus experiencing an increase in stability at the ice margin. These LLJ occurred, however, at times that did not match the theory.

3.3 Gusts

Wind gusts are typically generated by downward turbulent transport of momentum from higher altitudes (Suomi et al., 2013). Hence, in a tethersonde-based individual wind profile, a wind speed maximum at some layer may be simply due to a wind gust. This was studied by comparing the ascending and descending profiles (their time difference at the jet core height was never larger than 1 h). If a jet is only present in one of them, it suggests the influence of a gust. The data included four LLJs with the wind gust as the probable generation mechanism.

3.4 Fronts

We used the ECMWF operational analyses to detect fronts. During five soundings with a LLJ, a front was detected within a distance of about 800 km. All these LLJ cases were simultaneously related to either (a) baroclinicity, (b) inertial oscillations or (c) gusts. This is not surprising, as (a) non-occluded fronts are baroclinic, (b) in case of a cold front, the cold air mass typically penetrates below the warm air mass, building a stably stratified layer in between, which favours the generation of inertial oscillations, and (c) wind in the cold air mass is very often gusty (Wallace and Hobbs, 2006).

3.5 Summary of generation mechanisms

To summarize the potential forcing mechanisms, 13 cases suggested baroclinicity (3 of them also potentially inertial oscillation), four cases suggested non-baroclinic inertial oscillations, and four were related to wind gusts. Five of the above-mentioned cases were associated with frontal passages. The analysis of generation mechanisms explained 21 of the 43 LLJs observed. Accordingly there were 22 LLJs with an unexplained generation mechanism. 15 of these occurred during three intensive sounding

periods (each 10–17 h long), and seven occurred during other soundings. We note that seven of the unexplained LLJs occurred within 6 h from a case classified as baroclinic. Bearing in mind the uncertainties in the pressure analyses in the central Arctic (Inoue et al., 2009; Tetzlaff et al., 2012), this leaves open the possibility that also some of the unexplained LLJs were generated by baroclinicity.

4 Properties of low-level jets

To count the properties of LLJs, 25 cases were included (see Sect. 2.2). The jet core typically occurred at a height of 100–500 m (Fig. 5a), but the lowest one was observed at 70 m and highest at 1150 m altitude. On average, baroclinic jets were located lower and jets generated by gust higher than the others (Fig. 5a).

The most common depth of a jet was 400 to 600 m (Fig. 5b); only two sounding profiles showed a jet depth exceeding 1 km. Almost half of the jets with a depth exceeding 700 m were baroclinic (Fig. 5b). The average jet core wind speed (U_j) was 7.1 m s^{-1} (Fig. 5c; note that measurements were carried out only during winds lower than 15 m s^{-1}). The change in the wind speeds between the jet core and the subsequent minimum (strength of the jet) was less than 3.0 m s^{-1} in 44 % of the cases and stronger than 4.0 m s^{-1} in 28 % of the cases (Fig. 5d). Jets with the highest U_j were not the strongest ones. Many jets of unknown forcing mechanisms had a larger U_j than average (Fig. 5c) but a weaker jet: only in a single case a jet with an unknown origin was stronger than 4 m s^{-1} (Fig. 5d). If all soundings had reached the height of 2 km, there might have been some more cases of a stronger and deeper jet.

Considering all 25 jets, there was no correlation ($r = 0.004$) between the height of the jet core (z_j) and z_{Ri} (Fig. 6). However, the four soundings with inversion base temperatures (T_b) lower than −15 °C showed a high correlation ($r = 0.95$; $p = 0.052$). Sounding data showed that a jet core with higher than average wind speed ($U_j > 7.1$ m s^{-1}) occurred more often inside the turbulent layer (Fig. 6) (77 % of these cases showed $z_j < z_{Ri}$). Jet cores with smaller than average wind speed ($U_j < 7.1$ m s^{-1}) appeared above the turbulent layer (83 % of these cases showed $z_j > z_{Ri}$). A significant correlation ($r = 0.61$; $p < 0.01$) was observed between U_j and z_{Ri}. As the wind speed squared is a denominator in $Ri(z)$, the stronger is the wind, the smaller is $Ri(z)$. The smaller is $Ri(z)$, the higher is z_{Ri}. Among the LLJs with higher than average wind speed, 38 % were baroclinic, whereas only 25 % of LLJs with smaller than average wind speed were baroclinic (Fig. 5c).

Also the jet core height and the height of the top of temperature inversion z_t correlated ($r = 0.62$; $p < 0.01$; Fig. 7). LLJs with the core inside the turbulent layer had no significant correlation with z_t, LLJs with the core above the turbulent layer had a correlation coefficient of 0.72 ($p < 0.01$). The lower was the inversion base temperature (T_b), the higher was U_j ($r = -0.64$; $p < 0.01$).

5 Air mass origin

The 72-h backward trajectory calculations for the 25 LLJ cases observed (maximum one per day) showed that in most cases the air mass including a LLJ originated from the sea ice zone, with only seven cases from the open ocean. Even during these seven cases the air mass had traveled 800–1300 km over sea ice, as Tara was close to the North Pole.

The air mass origin (Fig. 8) was divided into five sectors: (1) 20° W–30° E (Fram Strait region), (2) 30°–165° E (Russian Arctic), (3) 165°–210° E (region towards the Bering Strait), (4) 210°–340° E (western Arctic), and (5) vicinity of the North Pole (northward of 85° N). All LLJs originating from the Fram Strait region (four cases) were located inside the turbulent layer ($z_j < z_{Ri}$) whereas all LLJs originating from the Russian Arctic (five cases) were located above the turbulent layer ($z_j > z_{Ri}$). There was not statistically significant difference between z_j of these groups, but the mean z_{Ri} of the Russian sector (174 m) was 171 m lower than the mean z_{Ri} of the Fram Strait sector ($p < 0.05$). In all cases of the Fram Strait sector, the air mass had been over the open sea less than 72 h before the LLJ was observed at Tara. In only one case of the Russian sector, the air

mass had been over the open sea (in this case z_j was only 20 m higher than z_{Ri}). The fifth sector included six cases and showed some differences from the other LLJs. The average jet depth of 803 m was as much as 356 m larger ($p < 0.05$) than in the case of other jets, and the mean z_a of 1390 m was 631 m higher ($p < 0.01$) than in the case of other LLJs.

6 Discussion and conclusion

The Tara tethersonde soundings probably represent the best data set of LLJs over the central Arctic Ocean from April through August (although late summer has been better covered by ship-based observations; Tjernström et al., 2012). The Tara results showed a lower occurrence of LLJs (46 %) compared to 80 % of Andreas et al. (2000) over the Antarctic sea ice. Also ReVelle and Nilsson (2008) suggested a high occurrence of LLJs (60–80 %) over polar oceans, and Vihma et al. (1998) observed that 91 % of rawinsonde soundings in the very baroclinic ice-edge zone in the Denmar Strait included a LLJ. According to our understanding, the most important reasons for the relatively low occurrence of LLJs at Tara were that (a) the observations were made far from strongly baroclinic zones, such as the sea ice margin, and (b) the typical conditions in April–August were not as stably stratified than in the autumn–winter data set of Andreas et al. (2000). Another data with a low occurrence of LLJs (25 %) were the year-round SHEBA rawinsonde soundings, which were taken far from the ice edge (Tjernström et al., 2004).

Jets with a high U_j occurred mostly inside the turbulent layer, and jets with a low U_j above the turbulent layer. Strong jet core winds contribute to growth of the turbulent layer, i.e. there is a two-way interaction between the ABL structure and LLJs. Previous studies have indicated some correlation between the jet core height z_j and the temperature inversion top height z_t ($r = 0.53$ in a climatology of LLJs over the USA, Bonner, 1968). We detected a more complex relationship: if the jet core was inside the turbulent layer, there was no significant correlation between z_j and z_t, whereas $r = 0.72$

($p < 0.01$) was observed for cases with the jet core above the turbulent layer. This is probably related to the common situation that in conditions of a strong temperature inversion, the turbulent layer is thin and inertial oscillations prevail, generating a jet close to z_t.

Our results for the typical jet core height (100–500 m) fit well with those of the Arctic Ocean Expedition 2001 (Tjernström, 2004), where the jet core typically occurred at the height of 200–400 m, whereas the core wind speeds were smaller (5–7 $m s^{-1}$) than in our data (7.1 $m s^{-1}$). The latter is somewhat surprising, as our data set was restricted to conditions of weak and moderate winds allowing tethersonde operation. A reason for the low winds in Tjernström (2004) might be related to the less good capability of rawinsonde soundings to detect jet cores. Also, similarly to the observations of Tjernström et al. (2004) and Andreas et al. (2000) over sea ice, we found the LLJ cores commonly within the temperature inversion layer (Fig. 7). In the observations of Vihma et al. (2011) over Svalbard fjords, LLJs were typically located above the top of the temperature inversion. These contrasting results were probably due to orographic effects.

According to Andreas et al. (2000), z_j and z_{Ri} agree very well. In our study, only the four cases with inversion base temperatures (T_b) under −15 °C showed a significant correlation ($r = 0.95$; $p = 0.05$). Our measurements were carried out in spring and summer, whereas those of Andreas et al. (2000) were taken in autumn and winter. In their data set, the temperature at the inversion base was most of the time less than −15 °C. Walter and Overland (1991) detected a LLJ during a research aircraft flight nearby the last sounding site of Tara. It is noteworthy that this cold-season LLJ was located at the top of the slightly stable layer just below the level where the Richardson number became very large, fitting very well to our population of cases with temperature less than −15 °C (Fig. 6).

The most important forcing mechanism for LLJs was baroclinicity, but also cases related to (potential) inertial oscillations, gust, and fronts were detected. The inertial oscillations were probably due to synoptic-scale changes in stratification. Our study

differed from many previous studies on LLJs in the sense that the diurnal cycle did not play a role, as the observations were made at latitudes 86–89° N.

Baroclinicity was responsible for generation of strong and warm LLJs, the former as in Vihma et al. (1998). Baroclinicity was a more important forcing mechanism in July and August (11 cases) than in April–June (two cases). The baroclinicity generating LLJs was mostly associated to transient cyclones, not to the climatological air temperature gradients. Accordingly, the July–August maximum may be related to the fact that in the central Arctic cyclones are more common (albeit weak) in summer than in any other season (Serreze and Barrett, 2008). In spring, the largest climatological temperature gradients occur over the sea ice margins, but Tara was far from these regions. Contrary to previous studies (Smedman et al., 2001), in the Tara data the baroclinic jets occurred at lower altitudes than other jets. This is probably because of the prevailing stable stratification in the Arctic, which keeps convection shallow and does not allow surface-generated horizontal temperature gradients to reach high altitudes.

In the case of approximately 50 % of the observed LLJs the generation mechanism remained unclear. Potential mechanisms that could not be detected by the tethersonde soundings and ECMWF analyses include mesoscale baroclinicity. Surface heating over areas of reduced sea ice concentration can generate horizontal temperature differences in the ABL (e.g. Vihma, 1995; Lüpkes et al., 2008). These are not necessarily reproduced by the ECMWF analyses, because the information on sea ice concentration is seldom accurate enough (Valkonen et al., 2008) and north of 84° N the sea ice concentration in the analyses was set to a constant value of 100 %, which was far from truth in summer 2007 (Comiso et al., 2008). If the sea ice zone includes large areas of open water that are not present in the ECMWF ice concentration field, a LLJ may also be generated via a spatial change in stratification and roughness (Vihma et al., 2003).

In addition, according to Kallistratova and Kouznetsov (2011), LLJs could be formed by the combined effects of baroclinicity, varying horizontal pressure gradient, and variations in the layered structure of inversions, which could lead to different amplitudes and initial phases of oscillations in different layers. In all cases that we were able to classify by the forcing mechanism, the temperature profile only included a single inversion layer. On the contrary, there were four cases with an unclear forcing mechanism in which the profile included two inversion layers. This may suggest the generation mechanism proposed by Kallistratova and Kouznetsov (2011).

We finally note that a forcing mechanism was found for almost all stronger LLJs and, as seen in Fig. 5c, d, the unclear cases mostly represented a weak jet strength under conditions of a strong wind. Further research needs on LLJs in the Arctic include investigation of low-level jet streams (Stensrud, 1996) and their effects on the sea ice drift and atmospheric moisture transport.

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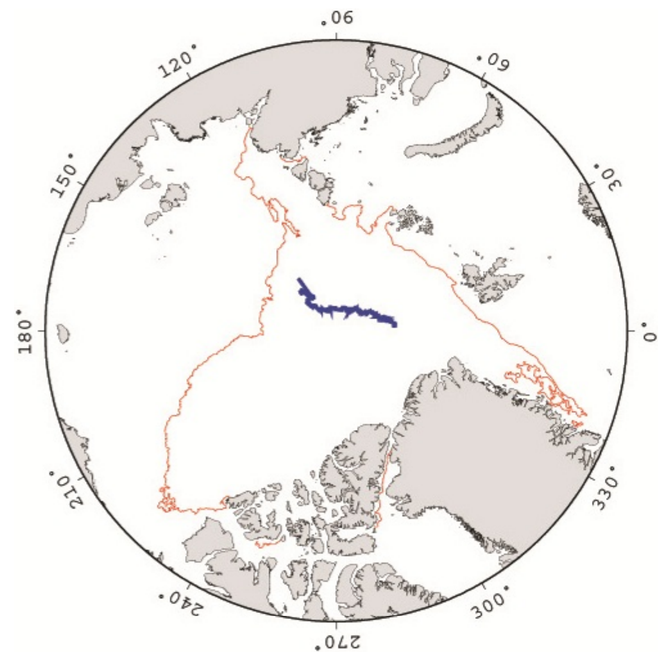

Fig. 1. Drift trajectory of Tara (blue) from the period of tethersonde soundings: 25 April to 31 August 2007. The brown line shows the September minimum sea ice extent.

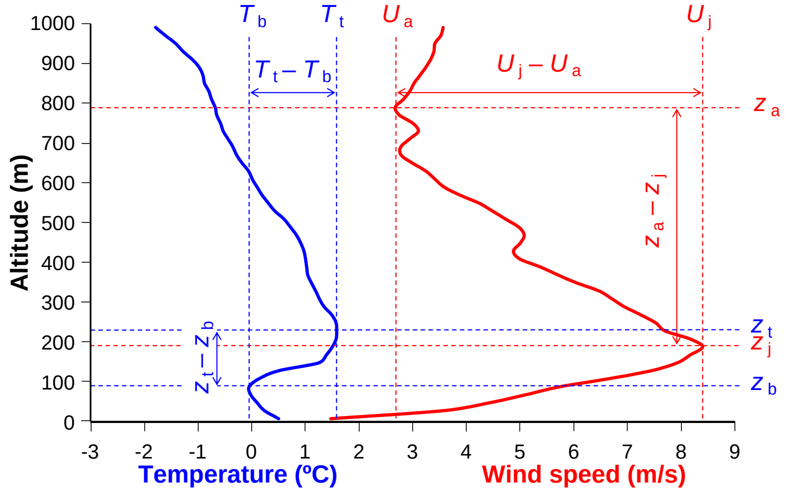

Fig. 2. Example of a tethersonde sounding at 13:00 UTC on 10 August 2007. The variables plotted are wind speed and temperature, provided as an illustration of the definitions used.

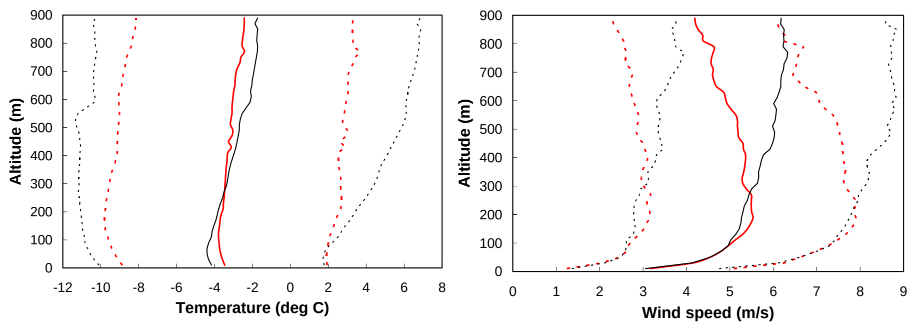

Fig. 3. The average profiles of **(a)** air temperature, and **(b)** wind speed based on all 43 profiles with LLJs observed (red) and on the 52 profiles without a LLJ (black). The dotted lines indicate the mean ± standard deviation.

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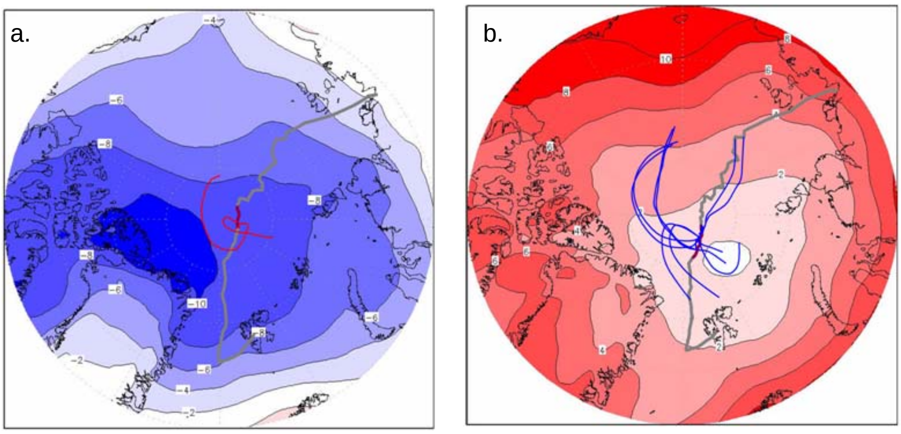

Fig. 4. The seasonal mean 925 hPa temperature field in **(a)** April–June and **(b)** July–August, based on the ECMWF operational analyses. The 72-h backward trajectories of baroclinic LLJs detected during these periods are marked by red (two cases in April–June) or blue curves (11 cases in July–August). The drift track of Tara is marked in gray with the sounding period highlighted in red.

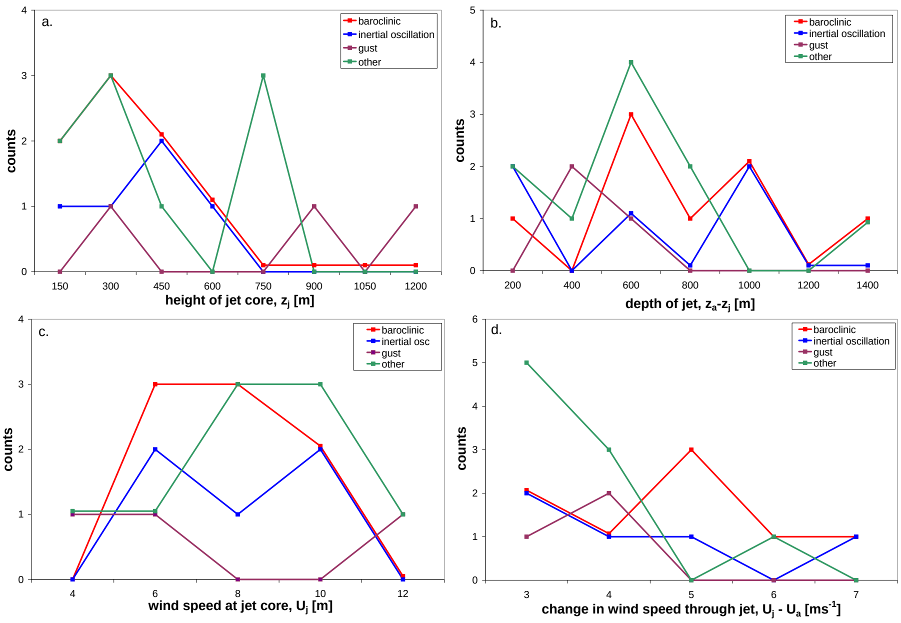

Fig. 5. Summary of the LLJ properties (indicated as counts): **(a)** the height of the jet core (z_j), **(b)** the depth of the jet ($z_a - z_j$), **(c)** the wind speed at the jet core (U_j), and **(d)** the change in wind speed trough the jet ($U_j - U_a$).

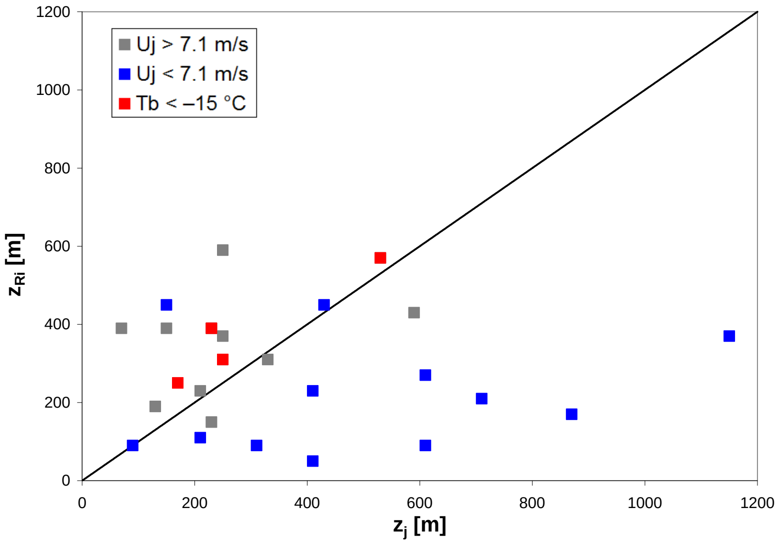

Fig. 6. A comparison of the height of the jet core, z_j, and height of the turbulent layer, z_{Ri}. LLJs are divided into groups of higher and lower than average U_j. Also four cases are shown where the inversion base temperature (T_b) is less than $-15\,^{\circ}$C (these cases belong to the higher U_j group).

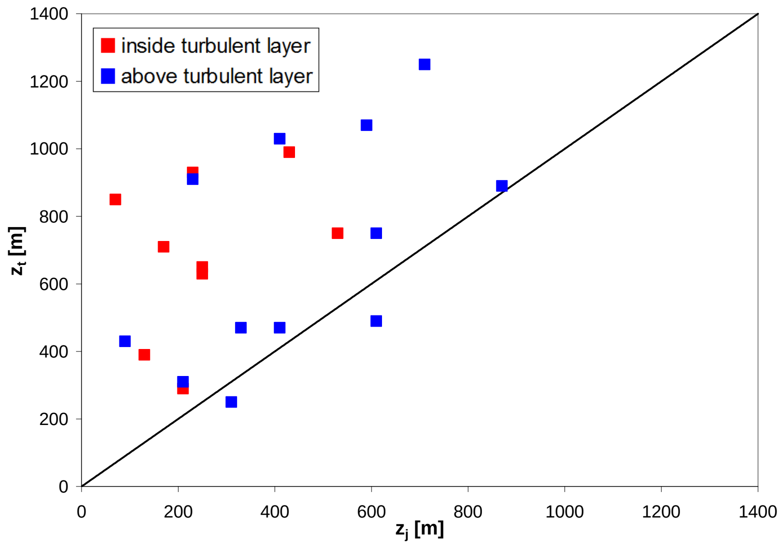

Fig. 7. A comparison of the height of the jet core, z_j, with the height of temperature inversion top, z_t. LLJs are divided into two groups with the core inside or outside the turbulent layer.

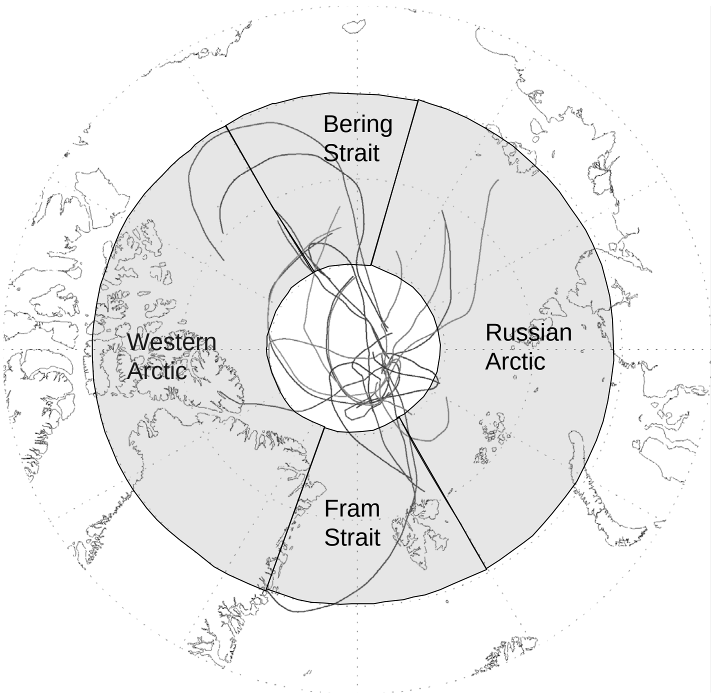

Fig. 8. 72-h backward trajectories (black curves) for 25 LLJs observed at Tara. The air mass origins are divided into sectors: Fram Strait, Russian Arctic, Bering Strait, Western Arctic and the sector in the vicinity of the North Pole (northward of 85° N).